## Diffusion of epicenter of earthquake aftershock, Omori's law, and generalized continuous-time random walk models [Helmstetter & Sornette, 2002b]

2017.5.29

### So Ozawa

(ERI, Hatano Lab, M1)

In these series of paper, authors derive many of empirical laws of earthquake by ETAS model.

- Sornette and Sornette, 1999
- Helmstetter and Sornette, 2002a
- Sornette and Helmstetter, 2002

In this paper [Helmstetter and Sornette, 2002b], we investigate **aftershock diffusion**.

## **Aftershock diffusion**

- from 1 km/h to 1 km/year
- Not universally observed





Mogi, 1968

### Why diffuse ?

- Viscous relaxation process (Rydelek and Sacks, 2001)
- Fluid transfer (Noir et al, 1997, Nur and Booker, 1972, Hudnut et al, 1989)
- Rate and State friction's law and non-uniform stress (Dieterich, 1994)

• Cascade process : Large aftershocks reproduce their secondary aftershocks close to them. (this paper)

## Flow

### 2. The ETAS model

Formulate ETAS model and refer the property of the model. Numerical simulation.

### 3. Mapping of the ETAS model on the CTRW model

Derive the master equation of ETAS.

Establish a correspondence between the ETAS model and the CTRW (Continuous Time Random Walk model).

#### 4. critical regime n=1

Derive the joint probability distribution N(t,r)

Calculate the average distance between mainshock and its aftershock R as a power law function of elapsed time.  $(R \sim t^{H})$ 

### 5. New Question on Aftershocks derived from the CTRW Analogy

### 6. Discussion

Summarize result of different regime Comparison to related study

### 7. Conclusion

## 2. The ETAS model

Formulate ETAS model and refer the property of the model. Numerical simulation.

## **ETAS Model**

$$\phi_{m_i}(t-t_i, \vec{r}-\vec{r}_i) = \rho(m_i)\Psi(t-t_i)\Phi(\vec{r}-\vec{r}_i).$$

 $m_i$  : magnitude  $r_i$  : positon  $t_i$  : time

'bare propagator' = seismic rate **directly** induced by a single 'mother' *i* 

(1) Large earthquake reproduce many aftershocks.

$$\rho(m_i) = K 10^{\alpha(m_i - m_0)},$$

(2) Normalized waiting time distribution = 'bare' omori's law  $\Psi(t) = \frac{\theta c^{\theta}}{(t+c)^{1+\theta}} H(t), \qquad \theta > 0, H(t) \text{ is Heaviside function}$ 

(3) Normalized spatial 'jump' distribution = isotropic elastic Green function dependence

$$\Phi(\vec{r}) = \frac{\mu}{d\left(\frac{|\vec{r}|}{d} + 1\right)^{1+\mu}}, \qquad \mu > 0$$

Seismogenesis Seminar

### $\alpha$ and b

event-size distribution = GR law number of daughter  $P(m) = b \ln(10) 10^{-b(m-m_0)}, \quad (6) \qquad \rho(m_i) = K 10^{\alpha(m_i - m_0)}, \quad (3)$ 

 $\alpha > b$ : large event dominate earthquake triggering  $\alpha < b$ : small event dominate earthquake triggering

recent reanalysis of seismic catalogs indicates  $\alpha < b$  and  $\alpha = 0.8$  (Helmstetter, 2003)

but case of  $\alpha > 0.5$  is difficult to analyze (infinite variance  $\rho(m)$ )

therefore our model uses  $b = 1, \alpha = 0.5$ 

### branching ratio n (Helmstetter & Sornette, 2002a)

*n* : *average* # of daughter created per mother event (summed by all possible magnitude)

$$n \equiv \int d\vec{r} \int_{t_i}^{+\infty} dt \int_{m_0}^{+\infty} dm_i P(m_i) \phi_{m_i}(t-t_i, \vec{r}-\vec{r}_i)$$
$$= \int_{m_0}^{+\infty} dm_i P(m_i) \rho(m_i) = \frac{Kb}{b-\alpha},$$

due to cascades of aftershocks, total # of event is larger by the factor  $1/(1-n) \sim 10$  $\rightarrow$  n is a branching parameter

> n < 1: subcritical regime (finally die out) n > 1: supercritical regime (exponentially increase) n = 1: critical regime (border between birth and death)



t<t\*, all regime behave identically

### Numerical simulation : method (Ogata, 1998 & 1999)

### initial condition

t=0 r=0 M7 event occur

### • algorithm

decide time of next event by nonstationary poisson process (8)  $\lambda(t) = \sum_{t_i \leq t} K 10^{\alpha(m_i - m_0)} \frac{\theta c^{\theta}}{(t - t_i + c)^{1 + \theta}},$ 

- $\rightarrow$  decide magnitude by GR law
- $\rightarrow$  select mother in all preceding events by (2)  $\phi_{m_i}(t-t_i, \vec{r}-\vec{r}_i) = \rho(m_i)\Psi(t-t_i)\Phi(\vec{r}-\vec{r}_i)$ .
- $\rightarrow \text{ decide location of new event by (5)} \quad \Phi(\vec{r}) = \frac{\mu}{d\left(\frac{|\vec{r}|}{d}+1\right)^{1+\mu}},$
- parameter set

 $\theta = 0.2, b = 1, \alpha = 0.5, n = 1, \mu = 1, m_0 = 0, d = 1$ km, c = 0.001day

## **Numerical simulation : Result**



considerable diffusion occurs

## [30,70] yrs : fractal distribution



correlation dimension D  $\sim$ 1.5 [0,70]yrs : D  $\sim$ 1.85 [7,70]yrs : D  $\sim$ 1.7 reported active fault system: D = [1.65:1.95]

# 3. Mapping of the ETAS model on the CTRW model

Derive the master equation of ETAS. Establish a correspondence between the ETAS model and the CTRW (Continuous Time Random Walk model).

## From direct Omori's law To renormalized Omori's law direct Omori law

$$\phi_{m_{i} \to m}(t-t_{i}, \vec{r} - \vec{r}_{i}) = \rho(m_{i} \to m) \Psi(t-t_{i}) \Phi(\vec{r} - \vec{r}_{i}),$$
  

$$\rho(m_{i} \to m) = n \ln(10)(b-\alpha) 10^{\alpha(m_{i} - m_{0})} 10^{-b(m-m_{0})}.$$
  
daughter(m, r, t)

#### renormalized Omori law



$$N_{m}(t,\vec{r}) = S(t,\vec{r},m) + \int \vec{d}r' \int_{m_{0}}^{\infty} dm' \qquad (17)$$

$$\times \int_{0}^{t} d\tau \phi_{m' \to m}(t - \tau,\vec{r} - \vec{r'}) N_{m'}(\tau,\vec{r'}).$$
assumption : daughter's magnitude is independent of its mother
$$(GR \text{ preserved all time. It is adequate only if } \alpha \leq b/2)$$

$$N_{m}(t,r) = P(m)N(t,r) \quad \text{for } t > 0$$
magnitude m vanishes
$$N(t,\vec{r}) = S_{M}(t,\vec{r}) + \int d\vec{r'} \int_{0}^{t} d\tau \phi(t - \tau,\vec{r} - \vec{r'}) N(\tau,\vec{r'}),$$

$$t > 0, \qquad (18)$$

$$S_{M}(t,\vec{r}) = \delta(r) \delta(t) \rho(M)/n,$$

Master Equation of ETAS = renormalized Omori's law  $N(t,r) = E[\lambda(t)\Phi(r)]$ : Expectation value(1<sup>st</sup> moment)

### Continuous time random walk model (Montroll & Weiss, 1965)

- generalization of naïve Random Walk model continuous distribution φ(r, t) of spatial step (jump length) and time step (wating time)
- master equation of CTRW is identical to ETAS

$$N(t,\vec{r}) = S_M(t,\vec{r}) + \int d\vec{r'} \int_0^t d\tau \phi(t-\tau,\vec{r}-\vec{r'})N(\tau,\vec{r'}),$$
  
$$t > 0, \qquad (18)$$

- A) N(t, r) : PDF for the random walker to *Just* arrive at r at t.
- B)  $S_M(t, r)$ : initial condition of random walk,
- C) integral on (18) denote superposition of all possible paths just having arrived at r at t, weighted by a transfer function  $\phi$
- Therefore we can borrow the deep knowledge of CTRW for the understanding Earthquake clustering.

• N and W

N(t,r): PDF of just arriving at position r at time t W(t,r): PDF of being at position r at time t

$$N(t,\vec{r}) = S_M(t,\vec{r}) + \int d\vec{r'} \int_0^t d\tau \phi(t-\tau,\vec{r}-\vec{r'}) N(\tau,\vec{r'}), \quad W(t,\vec{r}) = \int_0^t dt' \left[ 1 - \int_0^{t-t'} dt'' \Psi(t'') \right] N(t',\vec{r}).$$
(18)
(19)

• using Laplace-Fourier transform

$$\hat{N}(\beta,\vec{k}) = \frac{\hat{S}_{M}(\beta,\vec{k})}{1 - n\hat{\Psi}(\beta)\hat{\Phi}(\vec{k})}, \quad \hat{W}(\beta,\vec{k}) = \frac{1 - \hat{\Psi}(\beta)}{\beta}\hat{N}(\beta,\vec{k}).$$
(20)

• CTRW models *transport phenomena* in heterogeneous media. considering earthquake as *transport of stress* in heterogeneous crust, correspondence between ETAS and CTRW is natural ?

### summary : correspondence between ETAS and CTRW

TABLE I. Correspondence between the ETAS (epidemic-type aftershock sequence) and CTRW (continuous-time random walk) models. "PDF" stands for probability density function.

	ETAS	CTRW
$\overline{\Psi(t)}$	PDF for a "daughter" to be born at time $t$ from the mother that was born at time 0	PDF of waiting times
$\Phi(\vec{r})$	PDF for a daughter to be triggered at a distance $\vec{r}$ from its mother	PDF of jump sizes
т	Earthquake magnitude	Tag associated with each jump
$\rho(m)$	Number of daughters per mother of magnitude <i>m</i>	Local branching ratio
п	Average number of daughters created per mother summed over all possible magnitudes	Control parameter of the random walk survival (branching ratio)
n < 1	Subcritical aftershock regime	Subcritical "birth and death"
n = 1	Critical aftershock regime	The standard CTRW
<i>n</i> >1	Supercritical exponentially growing regime	Explosive regime of the "birth and death" CTRW
$N(t,\vec{r})$	Number of events of any possible magnitude at $\vec{r}$ at time t	PDF of just having arrived at $\vec{r}$ at time t
$W(t,\vec{r})$	PDF that an event at $\vec{r}$ has occurred at a time $t' \le t$ and that no event occurred anywhere from $t'$ to $t$	PDF of being at $\vec{r}$ at time t

## 4. critical regime n = 1

Derive the joint probability distribution N(t,r) Calculate the average distance between mainshock and its aftershock R as a power law function of elapsed time. (R~t^H)

## space : Fourier transform

$$\Phi(r) = \frac{\mu}{d(r/d+1)^{1+\mu}}$$
(5)

• for 
$$\mu > 2$$
,  $\langle r^2 \rangle = \sigma^2$  (finite)  
 $\widehat{\Phi}(k) = 1 - \sigma^2 k^2 + O(k^o)$  with  $o > 2$  (23)

• for  $0 < \mu \le 2$ ,  $\langle r^2 \rangle$  = infinite (so-called Levy-flight)

$$\widehat{\Phi}(k) = 1 - \sigma^{\mu} k^{\mu} + O(k^o) \text{ with } o > \mu \quad ^{(24)}$$

$$\sigma = \begin{cases} d[\Gamma(1-\mu)]^{1/\mu}, & 0 < \mu < 1\\ \frac{d\pi}{\mu\Gamma(\mu-1)\sin(\pi\mu/2)}, & 1 < \mu < 2. \end{cases}$$
(25)

## time : Laplace transform

$$\Psi(t) = \frac{\theta c^{\theta}}{(t+c)^{1+\theta}} H(t), \quad (4)$$

for  $\theta < 1$ ,

$$\hat{\Psi}(\beta) = 1 - (\beta c')^{\theta} + \mathcal{O}(\beta^{\omega}) \quad \text{with} \quad \omega \ge 1, \qquad (26)$$
$$c' = c \left( \Gamma(1-\theta) \right)^{\frac{1}{\theta}}$$

### for small $\beta$ and k,

$$\hat{N}(\beta,\vec{k}) = \frac{\hat{S}_M(\beta,\vec{k})}{1 - n\hat{\Psi}(\beta)\hat{\Phi}(\vec{k})}, \longrightarrow \hat{N}(\beta,\vec{k}) = \frac{\hat{S}_M(\beta,\vec{k})}{1 - n + n(\beta c')^{\theta} + n\sigma^{\mu}k^{\mu}}.$$
(27)

• case n=1  

$$\hat{N}(\beta,\vec{k}) = \hat{S}_M(\beta,\vec{k}) \frac{1}{(\beta c')^{\theta} + (\sigma k)^{\mu}} \xrightarrow{\text{Analyzed in detail below}} (51)$$

• case n<1  

$$\hat{N}(\beta,\vec{k}) = \frac{\hat{S}_M(\beta,\vec{k})}{1-n} \frac{1}{1+(\beta t^*)^{\theta}+(kr^*)^{\mu}}, r^* = \sigma \left(\frac{n}{1-n}\right)^{1/\mu}, t^* = c \left(\frac{n\Gamma(1-\theta)}{|1-n|}\right)^{1/\theta},$$

$$\begin{bmatrix} t < t^* \text{ and } r < r^* \longrightarrow \text{ Same expression as for n=1} \\ \text{otherwise} \longrightarrow \hat{N}(\beta,\vec{k}) \approx \frac{\hat{S}_M(\beta,\vec{k})}{1-n} \frac{1}{1+(\beta t^*)^{\theta}} \frac{1}{(1+(kr^*)^{\mu}}.$$
(31)

N can be factorized : No diffusion

### $heta > 1, \mu > 2$

 $\widehat{\Phi}(k) = 1 - \sigma^2 k^2 + O(k^o) \text{ with } o > 2, \quad \widehat{\Psi}(\beta) = 1 - (\beta c')^{\theta} + \mathcal{O}(\beta^{\omega}) \quad \text{ with } \omega \ge 1,$ 

$$\hat{N}(\beta,\vec{k}) = \frac{\hat{S}_M(\beta,\vec{k})}{1 - n\hat{\Psi}(\beta)\hat{\Phi}(\vec{k})}, \quad \longrightarrow \quad N(\beta,k) = S_M(\beta,k)\frac{1}{\beta c' + \sigma^2 k^2}$$

in real domain

$$N(t, \vec{r}) \propto \frac{1}{(Dt)^{d/2}} \exp[-(\vec{r})^2/Dt]$$
 where  $D = \sigma^2/c'$ ,  
(33)

 $R = \langle |\vec{r}|^2 \rangle^{1/2} \sim t^H$  with H=0.5 : standard diffusion

But  $\theta > 1$  is not appropriate case of  $\theta < 1$ ?

## heta < 1 , $\mu > 2$

From complicated calculation,

$$z = \frac{Dt^{\theta/2}}{|\vec{r}|} \quad (36)$$

for small 
$$z (r \gg Dt^{\theta}/2)$$
  $N(t, \vec{r}) = \frac{c'^{-\theta}}{2Dt^{1-(\theta/2)}} \sum_{k=0}^{\infty} \frac{(-1)^k z^k}{k! \Gamma[(1-k)\theta/2]}$  (40)

for large 
$$z (r \ll Dt^{\theta}/2)$$
  $N(t,r) \sim \frac{c'^{-\theta}}{Dt^{1-(\theta/2)}} \left(\frac{|\vec{r}|}{Dt^{\theta/2}}\right)^{(1-\theta)/(2-\theta)} \times \exp\left[-\left(1-\frac{\theta}{2}\right)\left(\frac{\theta}{2}\right)^{\theta/(2-\theta)} \left(\frac{|\vec{r}|}{Dt^{\theta/2}}\right)^{2/(2-\theta)}\right].$  (42)

N(t,r) cannot be factorized = **diffusion** 

 $R \sim t^H$  with H=  $\theta/2$  : subdiffusion

heta < 1 ,  $\mu > 2$ 



## Numerical simulation $\theta = 0.2$ , $\mu = 3$



$$\theta < 1, \mu \leq 2$$

$$\hat{N}(\beta, \vec{k}) = \hat{S}_{M}(\beta, \vec{k}) \frac{1}{(\beta c')^{\theta} + (\sigma k)^{\mu}} \cdot \hat{W}(\beta, \vec{k}) = \hat{S}_{M}(\beta, \vec{k}) \frac{(\beta)^{\theta - 1} c'^{\theta}}{(\beta c')^{\theta} + (\sigma k)^{\mu}} \cdot R \sim t^{H} \text{ with } H = \frac{\theta}{\mu} : \text{ superdiffusion or subdiffusion}$$

$$Dt^{\theta/2}$$

z expansion for small z and 1/z expansion for lagre z,  $z = \frac{Dt^{\theta/2}}{|\vec{r}|}$  (36)

for small 
$$z (r \gg Dt^{\theta}/2)$$
  $N(t,\vec{r}) = \frac{\sin\left(\frac{\pi\mu}{2}\right)}{\sigma c'\pi} \frac{\Gamma(1+\mu)}{\Gamma(2\theta)} \left(\frac{c'}{t}\right)^{1-2\theta} \left(\frac{\sigma}{|\vec{r}|}\right)^{1+\mu}$  (59)  
 $p = 1 - 2\theta$ 

for large 
$$z (r \ll Dt^{\theta}/2)$$
  $N(t, \vec{r}) = \frac{c^{-\theta}}{D\pi\mu t^{1-\theta+\theta/\mu}} \sum_{m=0}^{+\infty} (-1)^m \times \left[ \mu z^{1-\mu-m\mu} \frac{\Gamma(1-(m+1)\mu)\sin((m+1)\mu\pi/2)}{\Gamma(-m\theta)} + \frac{z^{-m}}{m!} \frac{\pi\cos(m\pi/2)}{\sin[(m+1)\pi/\mu]\Gamma(\theta-(m+1)\theta/\mu)} \right].$  (61)

## N(t,r) for large z can be further classified

$$N(t,r) \simeq \frac{\Gamma(1-2\mu)\sin(\pi\mu)\sin(\pi\theta)}{c'\sigma\pi^2} \frac{\Gamma(1+\theta)}{(r/\sigma)^{1-2\mu}} \frac{1}{(t/c')^{1+\theta}}$$
for  $\mu < 0.5$ ,  $p = 1+\theta$ 

$$N(t,r) \simeq \frac{c'^{-\theta}}{c'\sigma\mu\Gamma(\theta-\theta/\mu)\sin(\pi/\mu)} \frac{1}{(t/c')^{1-\theta+\theta/\mu}}$$

for 
$$0.5 < \mu < 2$$
. (62)  
$$p = 1 - \theta + \frac{\theta}{\mu}$$

θ=0.2, μ=0.2

θ=0.2, μ=0.9



## Numerical simulation $\theta=0.2, \mu=0.9$



 $R \sim t^H \text{ with } H=0.25 \text{ (predicted H is 0.22)}$ 



averaging over 500 sample

### Other distribution $\Psi(t)$ and $\Phi(r)$

•  $\Psi(t)$  and  $\Phi(r)$  are not power law  $\Psi(t) = \lambda e^{-\lambda t} \quad \Phi(\vec{r}) = L_{\mu}(|\vec{r}|)$   $R \sim t^{H}$  with  $H = \frac{1}{\mu}$ : superdiffusion at large time  $r \ll (\lambda t)^{\frac{1}{\mu}}, N(t,r) \sim 1/t^{\frac{1}{\mu}}$ 

Despite  $\Psi(t)$  is exponential distribution, local Omori's law  $p = 1/\mu$  is generated constant seismic rate for n=1

• nonseparable bare propagetor = microscopic diffusion process embodied

$$\phi_{m_{i}}(t-t_{i},\vec{r}-\vec{r}_{i}) = \rho(m_{i})\Psi(t-t_{i})\Phi(|\vec{r}-\vec{r}_{i}|/\sqrt{Dt}),$$

$$\Phi(|\vec{r}-\vec{r}_{i}|/\sqrt{Dt}) = \frac{1}{\sqrt{2Dt}}\exp(-|\vec{r}-\vec{r}_{i}|^{2}/Dt).$$

$$N(t,\vec{r}) \sim \frac{1}{t^{1-\theta}}\frac{1}{\sqrt{2\pi Dt}}\exp(-|\vec{r}|^{2}/Dt)$$

 $R \sim t^H$  with H = 0.5: standard diffusion

## 6. Discussion

Summarize result of different regime Comparison to related study

## **Diffusion exponent**





## **Comparison to related research**

- Noir et al., 1997 (1989 DobiEQ sequence) H =0.5 due to fluid transfer
- **Tajima & Kanamori, 1985** (subduction zone) logarithmic or H=0.1diffusion
- Shaw, 1993 (California)
   no diffusion and p~1 ← θ ~ 0, very small H ?
- Dieterich, 1994 (RSF law)

aftershock zone expand but not grow as power law.

• Marsan et al, 2000 (several catalogs)

 $H=0.2 \leftarrow apparent diffusion due to their analysis method (counting uncorrelated events )$ 

 Sotolongo-Costa et al., 2000 (microearthquakes in Spain) interpreted sequence of earthquakes as a random walk process

 different from this paper ( identify sequence as a single CTRW )

## 7. Conclusion

cascade of aftershock induce aftershock diffusion.

- correspondence between ETAS and CTRW
- different regimes of diffusion
- seismic diffusion occur and should be observed only when p <1 and t<t\*
- No anomalous stress diffusion is needed.